Macroscopic control parameter for avalanche models for bursty transport

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Similarity analysis is used to identify the control parameter $R_A$ for the subset of avalanching systems that can exhibit self-organized criticality (SOC). This parameter expresses the ratio of driving to dissipation. The transition to SOC, when the number of excited degrees of freedom is maximal, is found to occur when $R_A \rightarrow 0$. This is in the opposite sense to (Kolmogorov) turbulence, thus identifying a deep distinction between turbulence and SOC and suggesting an observable property that could distinguish them. A corollary of this similarity analysis is that SOC phenomenology, that is, power law scaling of avalanches, can persist for finite $R_A$ with the same $R_A \rightarrow 0$ exponent if the system supports a sufficiently large range of length scales, necessary for SOC to be a candidate for physical ($R_A$ finite) systems. © 2009 American Institute of Physics. [DOI: 10.1063/1.3057392]

I. INTRODUCTION

It is increasingly recognized that a large group of physical systems can be characterized as driven, dissipative, out-of-equilibrium, and having a conservation law or laws (see the comprehensive treatments of Refs. 1 and 2). They usually have many degrees of freedom (DOFs), or excited modes, and long range correlations leading to scaling or multiscaling. Two examples are fully developed turbulence (see, e.g., Refs. 3 and 4) and self-organized criticality (SOC5–7).

The SOC paradigm has found particular resonance with workers attempting to model, and understand, “bursty” scale free transport and energy release in magnetized plasmas (for a recent review, see, for example, Ref. 8). Simplified avalanche models have been proposed, and points of contact with the data are investigated, in the astrophysical context; most notably to describe magnetospheric activity (Refs. 9–12; see also Ref. 13 and references therein), the dynamics of the solar corona (e.g., Refs. 14–18; see also Ref. 19), and accretion disks (e.g., Refs. 20–22). In the context of magnetically confined laboratory plasmas there have been extensive efforts to construct avalanche models that make points of contact with the system under study and to establish signatures characteristic of SOC dynamics in experiments (e.g., Refs. 23–37). There have also been attempts to establish whether the signatures of SOC can emerge from magnetohydrodynamic (MHD) or reduced MHD models (e.g., in the solar coronal context38,39).

Since the original suggestion of Bak et al. in Ref. 40 that SOC “...could be considered a toy model of generalized turbulence,” there has been continuing debate on the possible relationship of turbulence to SOC.41–45 Similarities in the statistical signatures of turbulence, and systems in SOC, have been noted (see, e.g., Refs. 46 and 47). In particular, it has recently been argued in the context of astrophysical plasmas that SOC and turbulence are aspects of a single underlying physical process (see Refs. 48 and 49 and references therein). However, the extent to which SOC, as opposed to turbulence, uniquely captures the observed dynamics in magnetically confined laboratory plasmas (see Refs. 50–52) or is indeed consistent with it (see Refs. 53 and 54) has been brought into question. Key observables such as power law distributions of patches are not unique to SOC systems (for an example, see Ref. 55; see also the comprehensive discussion in Ref. 1).

Our focus here is then to establish the macroscopic similarities and differences between turbulence and SOC in the most general sense. A central idea in physics is that complex and otherwise intractable behavior may be quantified by a few measurable macroscopic control parameters. In fluid turbulence, the Reynolds number $R_E$ expresses the ratio of driving to dissipation and parametrizes the transition from laminar to turbulent flow. Control parameters such as the Reynolds number can be obtained from dimensional analysis (see, e.g., Refs. 3 and 56), without reference to the detailed dynamics. From this perspective the level of complexity resulting from the detailed dynamics is simply characterized by the number $N$ of excited, coupled DOF (or energy carrying modes) in the system. The transition from laminar to turbulent flow then corresponds to an (explosive) increase in $N$. The nature of this transition, the value of the $R_E$ at which it occurs, and the rate at which $N$ grows with $R_E$ all depend on specific system phenomenology. Dimensional arguments, along with the assumptions of steady state and energy conservation, are, however, sufficient to give the result that $N$ always grows with $R_E$ (as in Ref. 57; see also Ref. 3).

We anticipate that an analogous control parameter for complexity, $R_A$, will exist for the wider group of systems discussed above. Interestingly, it is now known that such a control parameter that expresses the ratio of driving to dissipation does indeed exist for SOC. In this paper we will give a prescription to obtain $R_A$ generally from dimensional analysis, that is, without reference to the range of detailed and rich phenomenology that any given system will also exhibit. The rate at which $N$ varies with $R_A$ is again dependent on this detailed phenomenology. We will see that similarity

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arguments, along with the assumptions of steady state and energy conservation, are, however, sufficient to determine whether or not $N$ grows with $R_k$.

The question of control parameters in SOC was initially controversial, as the name leads one to expect. It was originally argued (Refs. 5 and 40; see also Ref. 58) that avalanching systems self organized to the SOC state without a tuning parameter. Subsequent analysis has established a consensus (see Refs. 1, 2, and 59–62) that some tuning exists, at least in the sense that SOC is a limiting behavior in the driving rate $h$ and the dissipation rate $\epsilon$, such that $h/\epsilon \to 0$ with $h, \epsilon \to 0$ (and $h \leq \epsilon$, that is, a steady state). This understanding is exemplified in Jensen’s constructive definition as given in Ref. 7 of SOC as the behavior of “slowly driven interaction dominated thresholded” (SDIDT) systems. Clearly then, $h/\epsilon$ plays the role of a control parameter. This SDIDT limit $h/\epsilon \to 0$ has been investigated extensively (e.g., Refs. 60 and 61), most recently with respect to finite size scaling in the limit of increasingly large system size (e.g., Refs. 62 and 63).

Here we are concerned with the relevance of SOC to physically realizable systems, and, in particular, natural ones, where the system size is finite and the driving may be unknown and/or highly variable. Our focus is on parametrizing the level of complexity of the system as we take it away from the SDIDT limit by increasing the driver, in a system of large but fixed size. For avalanche models exhibiting SOC, we will argue that distinct realizable avalanche sizes play the role of excited DOF of the system. The SOC state is then characterized by maximal excited DOF, that is, avalanches occurring on all lengthscales supported by the system. Far from the SOC state, the system becomes ordered with few excited DOF and exhibits laminar flow. The SDIDT limit is reached by taking $R_k$ to zero, and we will show that this indeed maximizes the number of excited DOF $N$. The SDIDT limit is thus in the opposite sense to fluid turbulence which maximizes $N$ at $R_k \to \infty$.

This suggests a possible means to distinguish observationally between turbulence and SOC in observations and experiments of driven, magnetically confined plasmas. For example, a power law region in the power spectral density of some quantity that probes the flow is often identified in both laboratory and astrophysical confined magnetized plasmas (e.g., Refs. 27–30) and is discussed in the context of both SOC and turbulence. This power law region will always be of finite spatiotemporal range (an “inertial range” of the cascade). Our results imply that this inertial range will decrease as we increase the drive for SOC, whereas it will increase for turbulence—providing an experimental or observational test to distinguish these phenomena.

Our relationship between $R_k$ and $N$ implies the possibility of large but finite $N$ for small but nonzero $R_k$; hence an important corollary is that SOC phenomenology can quite generally persist under conditions of finite drive in a sufficiently large bandwidth system. This has been seen in specific avalanche models (see Refs. 64–66). Here, since our result flows from dimensional analysis, we will see that this is a generic property of avalanching systems.

II. SIMILARITY ANALYSIS AND CONTROL PARAMETER

We shall focus on how well-established techniques: similarity analysis (as described in Ref. 56) and the II theorem obtained by Buckingham in Ref. 67 nearly a century ago can be used to reveal new information about avalanche models exhibiting SOC (i.e., Refs. 1, 5–7, 58, and 59).

The systems that we have in mind all have strongly coupled excited DOF that transport some quantity from the driving to the dissipation scale. They have the following properties:

(I) The many excited DOFs of the system are coupled; there is some dynamical quantity that freely flows over all the excited DOFs of the system. We can characterize a flux $\epsilon_l$ of this quantity associated with processes that occur on lengthscale $l$, that is, $\epsilon_l$ is the transfer rate of the dynamical quantity through $l$ to neighboring lengthscales.

(II) The system is not necessarily in equilibrium but is in a steady state on the average.

(III) The dynamical quantity is conserved so that given (II) the injection rate $\epsilon_{\text{inj}}$ balances the dissipation rate $\epsilon_{\text{diss}}$, that is, $\epsilon_{\text{inj}} \sim \epsilon_{\text{diss}}$ in an ensemble averaged sense.

(IV) The solution is of a scaling type, that is,

$$N \sim \left( \frac{L_0}{\delta l} \right)^\alpha,$$

where $\alpha > 0$ and $L_0$ and $\delta l$ are the largest and smallest lengthscales, respectively, that are supported by the system.

(V) The number of excited DOF can be parametrized by a single macroscopic control parameter.

We will identify the control parameter for these systems in terms of known macroscopic variables by formal dimensional analysis (similarity analysis or Buckingham II theorem; see, e.g., Refs. 56 and 67). The essential idea is that the system’s behavior is captured by a general function $F$ which only depends on the relevant variables $Q_{1...V}$ that describe the system. Since $F$ is dimensionless it must be a function of the possible dimensionless groupings, the $\Pi_{1...M}(Q_{1...V})$, which can be formed from the $Q_{1...V}$. The (unknown) function $F(\Pi_1, \Pi_2, \ldots, \Pi_M)$ is universal, describing all systems that depend on the $Q_{1...V}$ through the $\Pi_{1...M}(Q_{1...V})$ and the relationships between them. If one then has additional information about the system, such as a conservation property, the $\Pi_{1...M}(Q_{1...V})$ can be related to each other to make $F$ explicit. Thus, this method can lead to information about the solution of a class of systems where the governing equations are unavailable or intractable, often the case for complex systems where there are a large number ($N$ here) of strongly coupled DOFs. If the $V$ macroscopic variables are expressed in $W$ physical dimensions (i.e., mass, length, and time) then there are $M=V-W$ dimensionless groupings.

The properties (I)–(V) above restrict the choice of relevant $Q_{1...V}$. First, we have only specified that there is a transfer rate on lengthscale $l$, $\epsilon_l$ [property (I)] of some dy-
TABLE I. II theorem applied to homogeneous turbulence.

<table>
<thead>
<tr>
<th>Variable</th>
<th>Dimension</th>
<th>Description</th>
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</thead>
<tbody>
<tr>
<td>$L_0$</td>
<td>L</td>
<td>Driving lengthscale</td>
</tr>
<tr>
<td>$\eta$</td>
<td>L</td>
<td>Dissipation lengthscale</td>
</tr>
<tr>
<td>$U$</td>
<td>LT$^{-1}$</td>
<td>Bulk (driving) flow speed</td>
</tr>
<tr>
<td>$\nu$</td>
<td>L$^2$T$^{-1}$</td>
<td>Viscosity</td>
</tr>
</tbody>
</table>

namical quantity, its precise nature is irrelevant. Consequently, the only physical dimensions of the transfer rate $\varepsilon$, relevant to the problem are length and time, so that $W=2$. Second, property (V) is that there is a single control parameter $\Pi_1$, which may be expressed as a function of the number of excited DOFs $N$. To incorporate the scaling property (IV) we will seek solutions such that $\Pi_1=g(L_0/\delta l)=f(N)$. This means that the system’s behavior is captured by some $F(\Pi_1, \Pi_2)=C$ which fixes $M=2$ ($C$ is a constant). The $\Pi_1$ and $\Pi_2$ are related to each other via properties II and III (conservation and steady state). We then have that $V=4$; there are always four relevant macroscopic variables to consider.

To see this in action, we begin with a relatively well understood example, namely, Kolmogorov (K41) turbulence. Our aim here is to straightforwardly illustrate the above approach by obtaining the control parameter, the Reynolds number $R_E$, as a function of $N$ via dimensional analysis; for a detailed discussion of the universal scaling properties of K41 turbulence and their origin in the Navier–Stokes equations see, for example, Ref. 3. As above, for K41 we have four relevant macroscopic variables (given in Table I) and two dimensionless groups,

$$\Pi_1 = \frac{UL_0}{\nu} = R_E, \quad \Pi_2 = \frac{L_0}{\eta} = f(N).$$

(2) $\Pi_1$ is just the Reynolds number $R_E$ of the flow, and the ratio of lengthscales $\Pi_2$ is related to the number of DOF $N$ that can be excited. We now see how $R_E$ is related to $f(N)$ by relating $\Pi_1$ to $\Pi_2$. For incompressible fluid turbulence, our dynamical quantity $\varepsilon$ is the time rate of energy transfer per unit mass through lengthscale $l$. The procedure is then as follows:

1. Conservation and steady state imply (ensemble averaged) that $\varepsilon_{\text{inj}} \sim \varepsilon_l \sim \varepsilon_{\text{dis}}$, that is, the average energy injection rate $\varepsilon_{\text{inj}}$ balances the average energy dissipation rate $\varepsilon_{\text{dis}}$.
2. The rate at which energy is transferred to the fluid is from dimensional analysis: $\varepsilon_{\text{inj}} \sim U^3/L_0$.
3. Dimensional analysis of Navier–Stokes gives $\varepsilon_{\text{dis}} \sim \nu^3/\eta^4$.
4. $\varepsilon_{\text{inj}} \sim \varepsilon_{\text{dis}}$ then relates $\Pi_1$ to $\Pi_2$,

$$R_E = \frac{UL_0}{\nu} \sim \left( \frac{L_0}{\eta} \right)^{\beta},$$

and fixes exponent $\beta=4/3$.

5. The solution is of scaling type, so that

$$N \sim \left( \frac{L_0}{\eta} \right)^{\alpha},$$

with $\alpha>0$ by definition.

(6) Thus

$$R_E \sim \left( \frac{L_0}{\eta} \right)^{\beta} \sim N^{\beta N},$$

(5) and $\beta_N=\beta/\alpha>0$.

The value of the exponents $\alpha$ and $\beta$ will depend on the detailed phenomenology of the turbulent flow. An estimate based on K41, for example, with $\beta=4/3$ from the above, and $\alpha=D=3$ where $D$ is Euclidean dimension, implies a high degree of disorganization and will be modified, for example, if the turbulence is intermittent. Importantly, the only property of turbulence with which we are concerned here is that both $\beta>0$ and $\alpha>0$ so that $\beta_N=\beta/\alpha>0$. This identifies the Reynolds number as the control parameter for a process (turbulence) which simply excites more active modes or DOF as we increase $R_E$.

We now see how the above arguments apply to other systems as defined above, in particular, to avalanche models. Without recourse to details of the system, similarity analysis will be sufficient to obtain the relationship between the control parameter $R$ and the number of DOFs $N$ of the form

$$R \sim N^{\beta N},$$

(6) The value of the exponent $\beta_N$ will depend on the details of these systems but crucially we will see that the sign of $\beta_N$ is fixed by the similarity analysis. This is sufficient to establish whether or not, as in the case of turbulence, increasing $R$ increases the number of excited DOFs in the system.

III. CONTROL PARAMETER FOR AVALANCHEING SYSTEMS

We now envisage a generic avalanche model in a system of size $L_0$ where the height of sand is specified on a grid, with nodes at spacing $\delta l$. Sand is added to individual nodes, that is, on lengthscales $\delta l$ at an average time rate $\varepsilon_{\text{inj}}=h$ per node. There is some process, here avalanches, which then transports this dynamical quantity (the sand) through structures on intermediate lengthscales $\delta l \ll L_0$. Sand is then lost to the system (dissipated) at a time rate $\varepsilon$ over the system size $L_0$. On intermediate lengthscales $\delta l \ll L_0$, sand is conservatively transported via avalanches (see also Refs. 68–70). In our discussion here we follow Ref. 5 and assume that the transport timescale is fast, so that avalanches occur instantaneously and do not overlap. There must be some detail of the internal evolution of the pile that maximizes the number of lengthscales $l$ on which avalanches can occur. For avalanche models this is the property that transport can only occur locally if some local critical gradient is exceeded; as a consequence the pile evolves through many metastable states. If these lengthscales represent excited DOF then the number $N$ of DOF available will be bounded by $L_0$ and $\delta l$ so that $N \sim (L_0/\delta l)^{\alpha}$, with $D\geq \alpha \geq 0$ for $D>1$ (if $\alpha$ may be fractional).
TABLE II. II theorem applied to an avalanching system. The sand carries a property with dimension $S$.

<table>
<thead>
<tr>
<th>Variable</th>
<th>Dimension</th>
<th>Description</th>
</tr>
</thead>
<tbody>
<tr>
<td>$L_0$</td>
<td>$L$</td>
<td>System size</td>
</tr>
<tr>
<td>$\delta l$</td>
<td>$L$</td>
<td>Grid size</td>
</tr>
<tr>
<td>$\epsilon$</td>
<td>$ST^{-1}$</td>
<td>System average dissipation/loss rate</td>
</tr>
<tr>
<td>$h$</td>
<td>$ST^{-1}$</td>
<td>Average driving rate per node</td>
</tr>
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</table>

The four relevant variables for the avalanching system are given in Table II. The two dimensionless groups are

$$\Pi_1 = \frac{h}{\epsilon} = R_A, \quad \Pi_2 = \frac{L_0}{\delta l} = f(N).$$

We will now relate the control parameter $\Pi_1 = h/\epsilon$ to the number of excited DOFs by following the same procedure as above. $\epsilon_1$ now refers to the rate of transfer of “sand” through lengthscale $l$.

1. Conservation and steady state imply (ensemble averaged) $e_{\text{inj}} \sim \epsilon_1 \sim e_{\text{diss}}$.
2. In Euclidean dimension $D$ there are $(L_0/\delta l)^D$ nodes; $D > 0$ by definition. The rate at which sand is transferred to the pile is then from dimensional analysis: $e_{\text{inj}} \sim h(L_0/\delta l)^D$.
3. The system average dissipation rate is defined as $\epsilon = e_{\text{diss}}$.
4. $e_{\text{inj}} \sim e_{\text{diss}}$ then gives $h(L_0/\delta l)^D \sim \epsilon$ or

$$R_A = \frac{h}{\epsilon} \sim \left( \frac{\delta l}{L_0} \right)^D,$$

thus in the above notation fixes $\beta = -D < 0$.
5. The number $N$ of DOF available will be bounded by $L_0$ and $\delta l$ so that

$$N \sim \left( \frac{L_0}{\delta l} \right)^\alpha,$$

with $D \geq \alpha \geq 0$ for $D > 1$ (the value of $\alpha$ depends on the details and may be fractional).
6. Thus

$$R_A = \frac{h}{\epsilon} \sim \left( \frac{\delta l}{L_0} \right)^D \sim N^{-\alpha D} \sim N^{\beta_N}$$

and $\beta_N = \beta/\alpha = -D/\alpha < 0$.

We then have that the number of excited DOF decreases as we increase the control parameter $R_A = h/\epsilon$. Thus we recover the SUSDIT limit for SOC, namely $R_A \sim 0$, but now explicitly identify this limit with maximizing the number of excited DOF. Our result from dimensional analysis is to obtain $R_A \sim N^{\beta_N}$ and to show quite generally that for the avalanching system $\beta_N < 0$.

Our dimensional analysis for the avalanche model maps onto that for K41 turbulence, so in that sense $R_A = R_E$, that is, $R_A$ is the avalanching system’s “effective Reynolds number,” which expresses the ratio of driving to dissipation. Both $R_E$ and $R_A$ increase with driving of the system, but the system’s response is quite different. In the case of K41 turbulence, the system can excite more modes or DOFs and the flow becomes more disorganized, whereas in the avalanche models, less DOFs are available so the system is pushed toward order. The essential difference between the two systems in this context is as follows. As we increase the driving in K41 turbulence, the smallest lengthscale $\eta$ can decrease (via Navier–Stokes) to provide the necessary dissipation to maintain a steady state, and since we have assumed scaling the system simply excites more modes or DOF. On the other hand, in the avalanche models both the smallest and largest lengthscales are fixed; increasing the driving will ultimately introduce sand at a rate that exceeds the rate at which sand can be transported by the smallest avalanches, as we discuss next.

IV. SOC-LIKE BEHAVIOR UNDER INTERMEDIATE DRIVE

For avalanching to be the dominant mode of transport of sand, there are conditions on the microscopic details of the system; specifically, there must be a separation of timescales such that the relaxation time for the avalanches must be short compared to the time taken for the driving to accumulate sufficient sand locally to trigger an avalanche. Avalanches are triggered when a critical value for the local gradient is exceeded, the critical gradient can be a random variable but provided it has a defined average value $g$, we have that on average, we would need to add $g \delta l$ sand to a single cell of an initially flat pile to trigger redistribution of sand. The number of timesteps that this would take to occur would on average be $(g \delta l)/(h \delta t)$ where again $\delta l$ is the cell size and $\delta t$ is the timestep. This gives the condition for avalanching to dominate transport on all lengthscales in the grid $[\delta l, L_0]$, so that avalanches only occur after many grains of sand have been added to any given cell in the pile and is the strict SUSDIT limit,

$$h \delta t \ll g \delta l.$$  

We will now consider an intermediate behavior (see also Ref. 64),

$$g \delta l < h \delta t \ll g \delta l \left( \frac{L_0}{\delta l} \right)^D,$$

where the driver is large enough to swamp of order $h \delta t/(g \delta l)$ cells of the pile at each timestep (each addition of sand), but this is still much smaller than the largest avalanches that the system is able to support since the largest possible avalanche in a system of Euclidean dimension $D$ is $(L_0/\delta l)^D$ cells.

For a given physical realization of the sandpile, that is, fixed box size $L_0$ and grid size $\delta l$, successively increasing $h \delta t$ above $g \delta l$ then successively increases the smallest avalanche size (to some $\delta l' > \delta l$). Ultimately as $h$ and hence $R_A$ is increased to the point where $h \delta t \sim g \delta l(L_0/\delta l)^D$ there will be a crossover to laminar flow, as each addition of sand drives avalanches that are on the size of the system.

We now assume that the avalanching process is self-similar, so that the system is large enough that the probability density of avalanche sizes $S$ is $P(S) \sim S^{-\gamma}$ over a large range of $S$; that is, finite size effects do not dominate. Consequently
this intermediate, finite $R_A$ behavior will be “SOC-like,” with avalanches occurring within the range of lengthscales $[\delta l', L_0]$ with power law statistics sharing the same exponent $\gamma$ as at the SDIDT limit.

We will illustrate these remarks with simulations of the Bak et al.\cite{bak1988self} sandpile in two dimensions (2D), where the driving occurs randomly in time and is spatially restricted to the “top” of the pile. In all cases shown, the critical gradient (threshold for avalanching) is $g = 4 \delta^{-1}$, and normalized distributions of the number of topplings in an avalanche $S$ are shown (we take topplings as a measure of avalanche size following Ref. \cite{bak1988self}). In Fig. 1 we plot the results from two simulations in a box of size $L_0/\delta l = 100$, under driving rates $h = 4 \delta^{-1}$ and $h' = 16 \delta^{-1}$. We can see that as we increase the driving rate from $h = 4 \delta^{-1}$ to $h' = 16 \delta^{-1}$, the occurrence probability of the smallest avalanches is reduced, and on these normalized histograms, the probability of larger events is increased. These larger events for the run with $h' = 16 \delta^{-1}$, that is, for $S \sim [10^2 - 10^{3.5}]$, still follow the same power law scaling as the $h = 4 \delta^{-1}$ run (the precise location on the plot of the crossover in behavior will depend on details of the dynamics of the pile). This is to be anticipated provided that transport on these intermediate scales is still dominated by avalanching, that is, intermediate scale avalanches still have the property that they relax on a timescale that is much faster than that required by the driving to initiate an avalanche. If this is the case, then the phenomenology of these intermediate scale avalanches is unchanged by the increase in the driving rate, and as a consequence, except close to the crossover in statistics, their scaling exponent is, as we see, unchanged. As the system has self-similar spatial scaling we can also anticipate obtaining the same solution for these avalanches subject to a rescaling; $S$, which is a measure of avalanche size, will simply scale with $h \delta t$, the sand which must be redistributed at each timestep since $h \delta t > g \delta l$. This is shown in the right hand plot of Fig. 1 where we have rescaled the $h' = 16 \delta^{-1}$ intermediate range driving results by $S \to S'/16$. We can see that power law regions of the plots that both correspond to avalanching now coincide.

We can go further and anticipate that two realizations of the system, one with $h$ and $L_0$ and the other with $h' = Ah$ and $L_0' = aL_0$, give the same solution for $P(S)$ under rescaling $S \to S'/A^D$. This is shown in Fig. 2 where we compare two runs of the sandpile (i) with $h = 4 \delta^{-1}$ and $L_0 = 100 \delta l$ and (ii) with $h' = 16$ and $L_0' = 400 \delta l$, in the same format as Fig. 1. We can indeed see a close correspondence of the avalanche statistics in the power law region of the plot once we have rescaled the $h' = 16 \delta^{-1}$ and $L_0' = 400 \delta l$ run by $S \to S'/16$ (at the largest $S$, the histograms do not precisely collapse under this self-affine scaling; see Ref. \cite{bak1988self} for a discussion of the finite size scaling properties of the model).

This establishes a general property of avalanching systems that has been seen in several representative SOC models, such as in Refs. \cite{bak1988self}. Depending on the details, specifically, provided that a separation of timescales for avalanching can be maintained, some SOC systems will show scaling in systems where the drive is, in fact, highly variable. One could argue that such robustness against fluctuations in the driving is necessary for SOC to provide a “working model” in real physical systems where the idealized SDIDT limit may not be realized.

**V. CONCLUSIONS**

We have used similarity and dimensional analysis to discuss high dimensional, driven, dissipating, out-of-equilibrium systems, in particular, avalanching systems that exhibit bursty transport that can be in a SOC state. These act
to transport a dynamical quantity (e.g., for the avalanche models, sand) from the driving to the dissipation scale, in a manner that is conservative, that is steady state on the average, and that shows scaling. The generic nature of this method of analysis implies that our results are not restricted to sandpile models per se, and have wider application to physical systems that show bursty transport and scaling. We have postulated that a “class” of these systems has a single control parameter $R$ which expresses the ratio of the driving to the dissipation and which can be related to the number of excited DOFs $N$. Dimensional analysis then leads to a relationship of the form $R \sim N^{\beta_N}$ and, without reference to any detailed phenomenology of the system, determines the sign of $\beta_N$.

We have focused on avalanche models that can exhibit SOC, for which the above identifies the control parameter $R_A=h/\epsilon$. The limit $R_A \to 0$ is just the well known S disorder limit of SOC. Specific avalanching systems will have different values of $\beta_N$ but will all share the essential property that we obtain here, that $\beta_N < 0$ so that $N$ is maximal under the limit of vanishing driving. Our formalism for SOC has close correspondence with that for Kolmogorov homogeneous isotropic turbulence. A minimalist interpretation of our results is that Kolmogorov turbulence maximizes the number of excited DOF $N$ under maximal (infinite) driving in contrast to SOC. A maximalist interpretation is that $R_A$ is analogous to the Reynolds number $R_e$. This establishes an essential distinction between turbulence and SOC. Practically speaking, it can, for example, arise because if we fix the outer, driving scale in Kolmogorov turbulence, the dissipation scale can simply adjust as we increase the driving. Since the system shows scaling, this acts to increase the available DOFs. Avalanching on the other hand is realized in a finite sized domain (box) and driven on a fixed, smallest scale, so increasing the driving beyond a certain point simply swamps the smallest spatial scales, thus reducing the available DOFs. Increasing the driving then pushes Kolmogorov turbulence toward increasingly disorganized flow and avalanching systems toward more ordered (laminar) flow.

A corollary is that SOC phenomenology, that is, power law scaling of avalanches, can persist for finite $R_A$ with the same exponent that is seen at the $R_A \to 0$ limit, provided the system supports a sufficiently large range of length scales. This has been seen previously for specific realizations of avalanche models, but is shown here to be quite generic and is a necessary property for SOC to be a candidate for physical $(R_A$ finite) systems. As the driving is increased, the excited number of DOFs (modes) decreases for SOC and increases for turbulence, so that in principle one could distinguish SOC from turbulence observationally by testing how the bandwidth (range of spatiotemporal scales), over which scaling is observed, varies with the driving rate.

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